

# Documentation for Gar6more2D

Julien Diaz and Abdelaaziz Ezziani

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This code compute the quasi-analytical solution of several wave equation in two layered media, using the Cagniard de Hoop method [1, 2, 9, 8, 7, 3, 5, 4]. It produces seismograms at given points.

The equations can be written in the general form.

$$A(y)\frac{\partial^2 U}{\partial t^2} - B(y)U = \delta(\mathbf{x} - \mathbf{x}_s) f(t), \quad x \in \mathbb{R}, y \in \mathbb{R} \quad (1)$$

where  $A$  and  $B$  are operators satisfying

$$\begin{aligned} A(y) &= A^+, B(y) = B^+, & y > 0, \\ A(y) &= A^-, B(y) = B^-, & y < 0. \end{aligned}$$

The code analytically compute the Green function  $u$  of the problem

$$A(y)\frac{\partial^2 u}{\partial t^2} - B(y)u = \delta(\mathbf{x} - \mathbf{x}_s) \delta(t), \quad x \in \mathbb{R}, y \in \mathbb{R} \quad (2)$$

and convolves it with the source function  $f$ . You can modify this function in the subroutine *lib/libgeneral/source.F90*. The convolution is done by a numerical integration, that is why the solution is only “quasi-analytical”. You can improve the accuracy of the solution by increasing the number of intervals used for the integration (the variable  $N_{int}$  in the data file *cdh2d.dat*)

## 1 Acoustic/acoustic

The code computes a seismogram at points  $(x_i, y)_{i=1, Nx}$  of the pressure solution of the equations

$$\frac{\partial^2 P^+}{\partial t^2} - c^{+2} \Delta P^+ = \delta(\mathbf{x} - \mathbf{x}_s) f(t), \quad x \in \mathbb{R}, y > 0$$

$$\frac{\partial^2 P^-}{\partial t^2} - c^{-2} \Delta P^- = \delta(\mathbf{x} - \mathbf{x}_s) f(t), \quad x \in \mathbb{R}, y < 0$$

We consider  $\mathbf{x}_s = (0, h)$  and we use the transmission conditions :

$$\left| \begin{aligned} P^+ &= P^-, \\ \rho^+ \frac{\partial P^-}{\partial y} &= \rho^- \frac{\partial P^+}{\partial y}. \end{aligned} \right.$$

on the interface  $y = 0$ . The code also computes the velocity given by the relation :

$$\frac{\partial \mathbf{V}^\pm}{\partial t} = -\frac{1}{\rho^\pm} \nabla P^\pm.$$

If you want to compute the displacement  $U$ , it can be easily computed by replacing  $f(t)$  by the primitive of the source function you are using. For instance, if you are using a Rickert, you'll have to consider a first derivative of a Gaussian for  $f$ .

## 2 Acoustic/elastodynamic(isotropic)

The code computes a seismogram at point  $(x_i, y)_{i=1, Nx}$  of the pressure (in the fluid) and the velocity (in the solid) solution of the equations

$$\frac{\partial^2 P^+}{\partial t^2} - c^{+2} \Delta P^+ = \delta(\mathbf{x} - \mathbf{x}_s) f(t), \quad x \in \mathbb{R}, y > 0, \quad (3)$$

$$\frac{\partial^2 \mathbf{V}^-}{\partial t^2} - (\lambda^- + 2\mu^-) \nabla(\nabla \cdot \mathbf{V}) + \mu^- \nabla \times (\nabla \times \mathbf{V}^-) = 0, \quad x \in \mathbb{R}, y < 0, \quad (4)$$

with  $\mathbf{x}_s = (0, h)$  and the transmission conditions

$$\left| \begin{array}{l} \frac{\partial V_y^-}{\partial t} = -\frac{1}{\rho^+} \frac{\partial P^+}{\partial y}, \quad y = 0, \\ (\lambda^- + 2\mu^-) \frac{\partial V_y^-}{\partial y} + \lambda^- \frac{\partial V_x^-}{\partial x} = \frac{\partial P^+}{\partial t} \\ \frac{\partial V_x^-}{\partial y} + \frac{\partial V_y^-}{\partial x} = 0. \end{array} \right.$$

on the interface  $y = 0$ . The code also computes the velocity in the fluid by using the relation

$$\frac{\partial V^+}{\partial t} = -\frac{1}{\rho^+} \nabla P^+.$$

Once again, if you want to compute the displacement  $U$ , it can be easily computed by replacing  $f(t)$  by the primitive of the source function you are using.

## 3 Acoustic/poroelastic(see [5])

The code computes a seismogram at point  $(x_i, y)_{i=1, Nx}$  of the potential of velocity  $\chi$  and the displacement  $U^+$  (in the fluid) and the solid displacement  $U_s^-$  (in the poroelastic medium) solution of the equations

$$\left| \begin{array}{l} \frac{\partial^2 \chi^+}{\partial t^2} - c^{+2} \Delta \chi^+ = \delta(\mathbf{x} - \mathbf{x}_s) f(t), \\ \frac{\partial U^+}{\partial t} = -\nabla \chi^+. \end{array} \right.$$

for  $x \in \mathbb{R}, y > 0$  and

$$\left| \begin{array}{l} (1 - \phi^-) \rho_s^- \frac{\partial^2 \mathbf{U}_s^-}{\partial t^2} + \phi \rho_f^- \frac{\partial^2 \mathbf{U}_f^-}{\partial t^2} - (\lambda^- + 2\mu^-) \nabla(\nabla \cdot \mathbf{U}_s^-) + \mu^- \nabla \times (\nabla \times \mathbf{U}_s^-) + \beta \nabla P^- = 0, \\ (1 - a^-) \rho_f^- \frac{\partial^2 \mathbf{U}_s^-}{\partial t^2} + a^- \rho_f^- \frac{\partial^2 \mathbf{U}_f^-}{\partial t^2} + \nabla P^- = 0 \\ \frac{1}{m^-} P^- + (\beta^- - \phi^-) \nabla \cdot \mathbf{U}_s^- + \phi^- \nabla \cdot \mathbf{U}_f^- = 0 \end{array} \right.$$

for  $x \in \mathbb{R}, y < 0$ , either with the open pore transmission conditions (if parameter open is set to 1)

$$\left| \begin{array}{l} \phi^- (U_{fy}^- - U_{sy}^-) = U_y^+ - U_{sy}^-, \\ P^- = \rho^+ \frac{\partial \chi^+}{\partial t}, \\ (\lambda^- + m^- \beta^- (\beta^- - \phi^-)) \nabla \cdot \mathbf{U}_s^- + 2\mu^- \frac{\partial U_{sy}^-}{\partial y} + m^- \beta^- \phi^- \nabla \cdot \mathbf{U}_f^- = -\rho^+ \frac{\partial \chi^+}{\partial t}, \\ \frac{\partial U_{sx}^-}{\partial y} + \frac{\partial U_{sy}^-}{\partial x} = 0, \end{array} \right.$$

or with the sealed pore transmission conditions (if parameter open is set to 0)

$$\left| \begin{array}{l} \phi^- (U_{fy}^- - U_{sy}^-) = U_y^+ - U_{sy}^-, \\ U_{fy}^- = U_{sy}^-, \\ (\lambda^- + m^- \beta^- (\beta^- - \phi^-)) \nabla \cdot \mathbf{U}_s^- + 2\mu^- \frac{\partial U_{sy}^-}{\partial y} + m^- \beta^- \phi^- \nabla \cdot \mathbf{U}_f^- = -\rho^+ \frac{\partial \chi^+}{\partial t}, \quad \text{on the inter-} \\ \frac{\partial U_{sx}^-}{\partial y} + \frac{\partial U_{sy}^-}{\partial x} = 0, \end{array} \right.$$

face  $y = 0$ . The code does not compute the fluid displacement and the pressure in the poroelastic medium, but there is no particular difficulty to do that.

**Remark 3.1** *The pressure in the fluid satisfies the relation*

$$P^+ = \rho^+ \frac{\partial \chi}{\partial t}.$$

*Therefore, if you want to compute a seismogram of the pressure and the velocities in the fluid and of the velocities in the solid, you have to replace  $f(t)$  by the derivative of the source function you are using. If you want to consider a source of pressure, you'll have to replace  $f(t)$  by the integral of the source function you are using.*

## 4 Poroelastic/poroelastic (see [4])

The code computes a seismogram at point  $(x_i, y)_{i=1, Nx}$  of the solid displacement  $U$  solution of the equations

$$\left| \begin{aligned} (1 - \phi^+) \rho_s^+ \frac{\partial^2 \mathbf{U}_s^+}{\partial t^2} + \phi^+ \rho_f^+ \frac{\partial^2 \mathbf{U}_f^+}{\partial t^2} - (\lambda^+ + 2\mu^+) \nabla(\nabla \cdot \mathbf{U}_s^+) + \mu^+ \nabla \times (\nabla \times \mathbf{U}_s^+) + \beta^+ \nabla P^+ &= \nabla \delta(\mathbf{x} - \mathbf{x}_s) F_s(t), \\ (1 - a^+) \rho_f^+ \frac{\partial^2 \mathbf{U}_s^+}{\partial t^2} + a^+ \rho_f^+ \frac{\partial^2 \mathbf{U}_f^+}{\partial t^2} + \nabla P^+ &= \nabla \delta(\mathbf{x} - \mathbf{x}_s) F_s(t) \\ \frac{1}{m^+} P^+ + (\beta^+ - \phi^+) \nabla \cdot \mathbf{U}_s^+ + \phi^+ \nabla \cdot \mathbf{U}_f^+ &= \delta(\mathbf{x} - \mathbf{x}_s) F_p(t) \end{aligned} \right|$$

for  $x \in \mathbb{R}, y > 0$  and

$$\left| \begin{aligned} (1 - \phi^-) \rho_s^- \frac{\partial^2 \mathbf{U}_s^-}{\partial t^2} + \phi^- \rho_f^- \frac{\partial^2 \mathbf{U}_f^-}{\partial t^2} - (\lambda^- + 2\mu^-) \nabla(\nabla \cdot \mathbf{U}_s^-) + \mu^- \nabla \times (\nabla \times \mathbf{U}_s^-) + \beta^- \nabla P^- &= 0, \\ (1 - a^-) \rho_f^- \frac{\partial^2 \mathbf{U}_s^-}{\partial t^2} + a^- \rho_f^- \frac{\partial^2 \mathbf{U}_f^-}{\partial t^2} + \nabla P^- &= 0 \\ \frac{1}{m^-} P^- + (\beta^- - \phi^-) \nabla \cdot \mathbf{U}_s^- + \phi^- \nabla \cdot \mathbf{U}_f^- &= 0 \end{aligned} \right|$$

for  $x \in \mathbb{R}, y < 0$ , with the transmission conditions on the interface  $y = 0$

$$\left| \begin{aligned} \phi^- (U_{fy}^- - U_{sy}^-) &= \phi^+ (U_{fy}^+ - U_{sy}^+), \\ \alpha^- \nabla \cdot \mathbf{U}_s^- + 2\mu^- \frac{\partial U_{sy}^-}{\partial y} + m^- \beta^- \phi^- \nabla \cdot \mathbf{U}_f^- &= \alpha^+ \nabla \cdot \mathbf{U}_s^+ + 2\mu^+ \frac{\partial U_{sy}^+}{\partial y} + m^+ \beta^+ \phi^+ \nabla \cdot \mathbf{U}_f^+, \\ \mu^- \left( \frac{\partial U_{sx}^-}{\partial y} + \frac{\partial U_{sy}^-}{\partial x} \right) &= \mu^+ \left( \frac{\partial U_{sx}^+}{\partial y} + \frac{\partial U_{sy}^+}{\partial x} \right), \\ U_{sx}^- = U_{sx}^+, \quad U_{sy}^- = U_{sy}^+, \quad P^- &= P^+, \end{aligned} \right|$$

with

$$\alpha^\pm = \lambda^\pm + m^\pm \beta^\pm (\beta^\pm - \phi^\pm).$$

Actually the code computes the solution for each source  $F_s$  and  $F_p$  separately. If you want a bulk source ( $F_s$ ), set the parameter *type\_source* to 1, if you want a pressure source ( $F_p$ ), set the parameter *type\_source* to 2.

**Remark 4.1** *The code does not really compute the displacement, but its derivative (for some reasons related to the Cagniard-de Hoop method, see [6, 4]). Therefore, you have to replace  $f(t)$  by the primitive of the source function you are using to compute the displacement.*

## References

- [1] L. Cagniard. *Reflection and refraction of progressive seismic waves*. McGraw-Hill, 1962.
- [2] A. T. de Hoop. The surface line source problem. *Appl. Sci. Res. B*, 8:349–356, 1959.

- [3] J. Diaz. *Approches analytiques et numériques de problèmes de transmission en propagation d'ondes en régime transitoire. Application au couplage fluide-structure et aux méthodes de couches parfaitement adaptées*. PhD thesis, Université Paris VI-Pierre et Marie Curie, 2005. in french.
- [4] J. Diaz and A. Ezziani. Analytical solution for wave propagation in stratified poroelastic medium. part I: the 2D case. Technical Report 6591, INRIA, 2008.
- [5] Julien Diaz and Abdelaâziz Ezziani. Analytical solution for waves propagation in heterogeneous acoustic/porous media. part i: the 2d case. *Communications in Computational Physics*, 7(1):171–194, 2010.
- [6] A. Ezziani. *Modélisation mathématique et numérique de la propagation d'ondes dans les milieux viscoélastiques et poroélastiques*. PhD thesis, Université Paris 9, 2005.
- [7] Q. Grimal. *Etude dans le domaine temporel de la propagation d'ondes élastiques en milieux stratifiés ; modélisation de la réponse du thorax à un impact*. PhD thesis, Université Paris 12-Val de Marne, 2003. in french.
- [8] Y. Pao and R. Gajewski. *The generalized ray theory and transient response of layered elastic solids*, volume 13 of *Physical Acoustics*, chapter 6, pages 183–265. 1977.
- [9] J. H. M. T. van der Hijden. *Propagation of transient elastic waves in stratified anisotropic media*, volume 32 of *North Holland Series in Applied Mathematics and Mechanics*. Elsevier Science Publishers, 1987.